

# Extracellular polymeric substances and algal bloom dynamics in sea ice

Anthony Jajeh<sup>1</sup>, Kenneth M. Golden<sup>2</sup>, Jody W. Deming<sup>3</sup>, Jody R. Reimer<sup>4</sup>

## Abstract

Algae living in the brine inclusions of polar sea ice produce a protective gelatinous coating of extracellular polymeric substances (EPS), which helps them survive in their harsh sea-ice habitat. One consequence of EPS production may be a reduction in nutrient availability, as EPS alters the ice microstructure, reducing permeability and the flux of nutrient-rich seawater into the ice. Here, we identify and analyze a novel biophysical feedback loop where algae modify their environment through EPS production, which then affects nutrient availability and subsequent algal growth. While mathematical models coupling nutrients and algae have been widely used to study bloom dynamics, here we introduce EPS as an additional variable to investigate this feedback and its role in microbial strategies for living in such an extreme environment as sea ice. Analysis of both transient and asymptotic dynamics suggests that EPS-mediated feedbacks suppress algal biomass in both the short and long term. Quasi-steady-state analysis of our model reveals a transcritical bifurcation governed by nutrient fluxes, indicating a threshold beyond which the system transitions from a barren state to one supporting microbial life. We also identify a Hopf bifurcation, which suggests that EPS-mediated negative feedbacks can give rise to sustained oscillations in nutrient and algal biomass levels. These results highlight how the tight coupling between EPS production, nutrient transport, and algal dynamics fundamentally shapes both sea-ice primary productivity and permeability.

## 1 Introduction

Polar sea ice hosts a dynamic community of microscopic inhabitants within its porous microstructure (Thomas and Dieckmann (2010)), which forms as seawater freezes. While much of the salt in seawater is expelled during freezing, some remains trapped within the ice, concentrated in liquid brine, creating a complex network of brine-filled inclusions (Lake and Lewis (1970); Petrich and Eicken (2009); Zong (2022)). The volume fraction, geometry, and connectivity of these inclusions depend strongly on temperature, crystalline structure, and conditions under which the ice was formed (Petrich and Eicken (2009); Thomas and Dieckmann (2010)). Ice algae dominate these brine inclusions in spring, when light becomes sufficient to support algal blooms in the ice (Stoecker et al. (2000)). Ice algae contribute significantly (up to 26% (Gradinger (2009))) to total primary production in the polar regions (Arrigo et al. (2008)). Due to their early availability in spring and lipid-rich composition (Gradinger (2009)), ice algae are foundational to the polar marine food web, supporting grazers such as zooplankton, which in turn support higher trophic levels, including fish, seabirds, whales, seals, and polar bears (Stoecker et al. (2000)). For example, a high percentage of a polar bear's diet (on average about 86% (Brown et al. (2018))) can be traced back to sea-ice algae.

Ice algae must endure extreme environmental conditions within the ice, including low temperatures, high salinity, and low light availability for much of the year. Many sea-ice microbes, including ice algae, have adapted by producing a cryoprotective gelatinous coating of extracellular polymeric substances (EPS), a complex mix of polysaccharides, proteins, and other components (Ewert and Deming (2013); Krembs et al. (2011)). In addition to acting as a buffer between the microbes and

---

<sup>1</sup>University of Utah, Department of Mathematics, Salt Lake City, UT, USA, [jajeh@math.utah.edu](mailto:jajeh@math.utah.edu) ORCID: 0009-0009-5546-5781

<sup>2</sup>University of Utah, Department of Mathematics and adjunct in Department of Biomedical Engineering, Salt Lake City, UT, USA. [ken.golden@utah.edu](mailto:ken.golden@utah.edu), ORCID: 0000-0002-1688-183X

<sup>3</sup>University of Washington, School of Oceanography, Seattle, WA, USA. [jdeming@uw.edu](mailto:jdeming@uw.edu), ORCID: 0000-0002-0310-1101

<sup>4</sup>University of Utah, Department of Mathematics and School of Biological Sciences, Salt Lake City, UT, USA. [jody.reimer@utah.edu](mailto:jody.reimer@utah.edu), ORCID: 0000-0001-7742-2728

40 their environment, EPS alters the physical properties of the surrounding ice (Krembs et al. (2011);  
 41 Roukaerts et al. (2021)). The presence of EPS, whether coating cells or released into their surround-  
 42 ings, changes the geometry of the pores and may impede fluid flow through clogging effects, altering  
 43 the flux of nutrients available to the microbial community (Krembs et al. (2011); Steffen et al.  
 44 (2018)). Reduced nutrient fluxes (in particular, nitrogen) limit algal growth (Gradinger (2009)), and  
 45 because algae are the largest producers of EPS, this is hypothesized to subsequently decrease EPS  
 46 production rates. This suggests a complex EPS-mediated biophysical feedback, with implications for  
 47 both sea-ice physics and ecosystem dynamics. For example, can a bloom be so successful with so  
 48 much EPS produced that the clogging, which can restrict vital nutrient fluxes, provides a check or  
 49 negative feedback on the growth, thus regulating the bloom?

50 In this paper, we take the first step toward understanding the biophysical feedback dynamics  
 51 mediated by EPS. We begin with a canonical algal bloom model and extend it to include EPS as  
 52 a new variable (Section 2). This extended model is a coupled set of ordinary differential equations  
 53 that describes the interactions among nutrients, algae, and EPS. While we do not explicitly model  
 54 sea ice microstructure, the effect of EPS on ice permeability is included through nonlinear terms  
 55 that capture its impact on nutrient fluxes. In Section 2.1, we analyze the most general form of  
 56 the model to understand its key features of asymptotic behavior. In Section 2.2, we explore model  
 57 sensitivity to the timescales over which EPS accumulate and decay, as these processes are difficult  
 58 to measure empirically and thus poorly constrained in our model. We analyze three model regimes  
 59 corresponding to three timescale assumptions: EPS dynamics are much slower than algal and nutrient  
 60 dynamics (Regime 1, Section 2.2.1); EPS dynamics are fast, responding instantaneously to changes  
 61 in algal biomass (Regime 2, Section 2.2.2); and EPS dynamics occur over similar timescales to those  
 62 of algae and nutrients (Regime 3, Section 2.2.3). We conduct a bifurcation and stability analysis  
 63 of our model under each of these assumptions. Comparison between regimes (Section 2.3) reveals  
 64 a persistent transcritical bifurcation that depends on nutrient flux terms, and a Hopf bifurcation  
 65 that is only possible if EPS dynamics occur over similar timescales to those of algae and nutrients  
 66 (Regime 3). This Hopf bifurcation suggests that strong EPS-mediated negative feedbacks may induce  
 67 periodic algal growth and, fascinatingly, concurrent periodic changes in ice permeability. Finally, in  
 68 Section 2.4, we investigate the system’s transient bloom dynamics through numerical simulations.  
 69 These simulations, in showing an inverse relationship between EPS accumulation and peak bloom  
 70 intensity, suggest that the negative EPS-mediated biophysical feedbacks that are the focus of this  
 71 paper do indeed reduce peak algal concentrations. Our model thus begins to address the questions  
 72 raised above concerning the role of EPS in regulating bloom dynamics.

## 73 2 Model formulation

74 We begin with a classic model of nutrient-driven phytoplankton blooms (Huppert et al. (2002)),  
 75 which captures the main processes and functional forms commonly found in sea ice biogeochemical  
 76 models (Lannuzel et al. (2020); Miller et al. (2015); Roukaerts et al. (2021); Tedesco and Vichi  
 77 (2014)).

$$\begin{aligned} \frac{dN}{dt} &= \alpha - \frac{\nu NA}{\gamma + N} - \beta N \\ \frac{dA}{dt} &= \xi \frac{\nu NA}{\gamma + N} - \delta A \end{aligned}$$

78 Typically, this model describes the dynamics of a limiting nutrient  $N$  and phytoplankton  $P$ ; however,  
 79 since ice algae are our primary producers of interest, we have changed the notation from  $P$  to  $A$ . We  
 80 have chosen nitrogen as the nutrient of interest, as nitrogen limitation is considered to inhibit ice algal  
 81 blooms (Kim et al. (2020)). Field measurements are often technically of particulate nitrogen, which  
 82 refers to nitrogen-containing compounds (both organic and inorganic) suspended in meltwater from  
 83 cores (Krembs et al. (2011)). Nutrient fluxes are described by input ( $\alpha$ ) and loss ( $\beta$ ) rates, assuming  
 84 nutrients enter the brine inclusions from the ocean below, facilitated by convective processes within  
 85 the ice (Petrich and Eicken (2009)). Nutrient loss by algae is described by the Monod term in both  
 86 equations, where  $\nu$  is the maximum uptake rate,  $\gamma$  is the half-saturation constant, and  $\xi$  is the rate of  
 87 algal growth corresponding to nutrient uptake. Algae are lost at a rate  $\delta$  due to death or loss from the  
 88 ice. All parameter values are assumed to be positive. Note that this model assumes constant, optimal  
 89 light for photosynthesis, and an unchanging ice environment corresponding to stable temperatures.

90 We now adapt this model to study EPS-mediated biophysical feedbacks. We add  $E$ , a state  
 91 variable describing the concentration of EPS in the system, often in units of milligrams of Xanthan

92 Gum equivalent per liter (mg XG/L) (Krembs et al. (2002)). We assume that the EPS accumulation  
 93 rate ( $\rho$ ) is proportional to the amount of algae present, and that it decays at a constant rate ( $\eta$ ).  
 94 As EPS accumulates, it clogs the brine inclusions, reducing nutrient fluxes (Krembs et al. (2011)).  
 95 To capture this process, we replace the parameters  $\alpha$  and  $\beta$  in the above equations with EPS-  
 96 dependent functions,  $\alpha(E)$  and  $\beta(E)$ . In general, these functions are assumed to be nonnegative and  
 97 monotonically decaying with  $E$ . The resulting model of coupled nutrient, algae, and EPS dynamics is

$$\begin{aligned}\frac{dN}{dt} &= \alpha(E) - \frac{\nu NA}{\gamma + N} - \beta(E)N \\ \frac{dA}{dt} &= \xi \frac{\nu NA}{\gamma + N} - \delta A \\ \frac{dE}{dt} &= \rho A - \eta E\end{aligned}\tag{1}$$

98 Table 1 describes the units of each model component and the values of model parameters, where  
 99 available, from the literature.

## 100 2.1 A nutrient-driven transcritical bifurcation

101 Empirical estimates of the best form for the nutrient flux functions,  $\alpha(E)$  and  $\beta(E)$ , do not exist, so  
 102 we first conduct a preliminary analysis of model (1) without explicitly specifying functional forms.  
 103 Setting  $\frac{dN}{dt} = \frac{dA}{dt} = \frac{dE}{dt} = 0$ , we find two steady-state solutions for  $(N^*, A^*, E^*)$ :

$$\begin{aligned}s_0 &:= \left( \frac{\alpha(0)}{\beta(0)}, 0, 0 \right) \\ s_1 &:= \left( \frac{-\delta\gamma}{-\xi\nu + \delta}, \frac{\xi\beta(E)}{\delta(\delta - \xi\nu)} \left( \gamma\delta + \frac{\alpha(E)}{\beta(E)}(\delta - \xi - \nu) \right), \frac{\rho A^*}{\eta} \right)\end{aligned}$$

104 Observe that  $s_0$  is a trivial steady state, characterized by the absence of both algae and EPS, while  
 105  $s_1$  is a nontrivial steady state in which algae and EPS persist. To assess the stability of  $s_0$ , we analyze  
 106 the Jacobian of (1) evaluated at this steady state (see S1), which yields the following eigenvalues:

$$\lambda_1 = -\eta, \quad \lambda_2 = -\beta(0), \quad \lambda_3 = \frac{\xi\nu\alpha(0) - \beta(0)\gamma\delta - \alpha(0)\delta}{\gamma\beta(0) + \alpha(0)}$$

107 Because all model parameters and  $\alpha$  and  $\beta$  are positive,  $\lambda_1$  and  $\lambda_2$  are negative, and the sign of  $\lambda_3$   
 108 depends on the sign of its numerator. Rearranging, we see that if  $\alpha(0)/\beta(0) < \gamma\delta/(\xi\nu - \delta)$ , then  
 109  $\lambda_3 < 0$  and  $s_0$  is a stable node. So if the ratio of physical nutrient input to loss in the absence of  
 110 EPS,  $\alpha(0)/\beta(0)$ , is sufficiently small, the system will eventually approach an algae- and EPS-free  
 111 state. However, if  $\alpha(0)/\beta(0)$  is sufficiently large, then  $\lambda_3 > 0$  and  $s_0$  is unstable. This scenario is a  
 112 transcritical bifurcation in the ratio of nutrient input to loss in an EPS-free system,  $\alpha(0)/\beta(0)$ . This  
 113 behavior provides a key ecological insight: algae persist only when the advective nutrient inflow is  
 114 sufficiently large relative to the physical nutrient outflow, replenishing the nutrients needed to sustain  
 115 the algal population. We will see this insight reflected in several places in the analysis that follows.

116 The nontrivial equilibrium  $s_1$  is only defined if all state variables are nonnegative, that is, if  
 117  $\frac{\alpha(E)}{\beta(E)} \geq \frac{\gamma\delta}{\nu\xi - \delta}$ , with  $\delta < \xi\nu$ . We cannot further analyze the existence and stability of  $s_1$  without  
 118 specifying functional forms for  $\alpha(E)$  and  $\beta(E)$ . Here we choose to use a negative exponential function,  
 119 which captures the hypothesis that higher EPS reduces nutrient fluxes. We define  $\alpha(E) = \phi \exp(-\frac{E}{\mu})$   
 120 and  $\beta(E) = \psi \exp(-\frac{E}{\mu})$ , where  $\phi$  and  $\psi$  represent the nutrient input and loss rates in the absence  
 121 of EPS, respectively, and  $\mu$  is the strength of EPS limitation. This approach results in the fully  
 122 specified model

$$\begin{aligned}\frac{dN}{dt} &= \phi e^{-\frac{E}{\mu}} - \frac{\nu NA}{\gamma + N} - \psi e^{-\frac{E}{\mu}} N \\ \frac{dA}{dt} &= \xi \frac{\nu AN}{\gamma + N} - \delta A \\ \frac{dE}{dt} &= \rho A - \eta E\end{aligned}\tag{2}$$

Symbol	Biological Description	Value	Range	Units	Source
$N$	Average concentration of particulate nitrogen	N/A	0.1–0.2	mg N/L	(Krembs et al. (2011)) Supp. Table S2
$A$	Mean chl-a in bottom 10 cm of sea ice	N/A	0.03–0.06	mg/L	(Krembs et al. (2011)) Supp. Table S2
$E$	EPS concentration	N/A	0.8–7.7	mg XG/L	(Krembs et al. (2002))
$N_0$	Initial condition for nutrient	0.2	-	mg N/L	Prescribed
$A_0$	Initial condition for algae	0.0002	-	mg A/L	Prescribed
$E_0$	Initial condition for EPS	0.002	-	mg XG/L	Prescribed
$\phi$	Inflow rate of nutrients	0.01	-	mg N/(L day)	Prescribed
$\psi$	Outflow rate of nutrients	0.01	-	1/day	Prescribed
$\nu$	Nutrient uptake rate by bottom ice algae	0.2	0.24	mg N/(mg chl-a day)	(Kim et al. (2020))
$\rho$	Production rate of EPS	.75	0.1–1	mg XG/(mg chl-a L day)	(Krembs et al. (2002))
$\delta$	Mortality rate of ice algae	0.007	0.007	1/day	(Hunke et al. (2024))
$\eta$	Bacterial degradation of polysaccharides	0.03	0.03	1/day	(Hunke et al. (2024))
$\gamma$	Half-saturation constant of ice algae nitrogen uptake rate	0.01	0.01	mg N/L	(Hunke et al. (2024))
$\xi$	Algal chl-a to nitrogen conversion ratio	0.2	0.2	mg chl-a/mg N	(Hunke et al. (2024))
$\mu$	Strength of EPS feedback	0.001	-	mg XG/L	Prescribed
$\sigma$	Scaling of how EPS tracks algae in Regime 2	0.05	-	-	Prescribed

**Table 1:** State variables and parameter values used throughout this work. Rows for the state variables  $N$ ,  $A$ , and  $E$  describe their observed average concentrations in spring sea ice for comparison with our model outputs. The Value column contains the parameter values chosen as the defaults for all figures produced in the text unless noted otherwise. The Range column contains the parameter ranges documented in the literature described under Source.

## 123 2.2 Model analysis under different timescale assumptions

124 Quantifying EPS in sea ice is difficult. While EPS can be detected and visualized in sea ice via Alcian  
125 Blue staining, the existing methods for quantitative measurements require melting the ice to extract  
126 the EPS, making in situ serial measurements currently impossible (Miller et al. (2015)). Because of  
127 this, little is known about how rapidly EPS accumulate or decay within sea ice. We explore three  
128 possible temporal regimes by modifying and analyzing (2) accordingly.

129 **2.2.1 Regime 1: EPS dynamics are much slower than algal and nutrient**  
 130 **dynamics**

131 We start by nondimensionalizing (2) (see Appendix 8) with this temporal scaling in mind, which  
 132 results in two fast equations and one slow one,

$$\text{Fast} \begin{cases} \epsilon \frac{dN}{d\tau} = ae^{-E} - \frac{cAN}{N+1} - bNe^{-E} \\ \epsilon \frac{dA}{d\tau} = \frac{fNA}{N+1} - A \end{cases} \quad (3a)$$

$$\text{Slow} \begin{cases} \frac{dE}{d\tau} = dA - E \end{cases} \quad (3b)$$

133 We now perform a quasi-steady-state analysis of our system. To analyze the fast set of equations,  
 134 we scale time by  $\epsilon$ , that is,  $\tau' = \frac{\tau}{\epsilon}$ . Then (3) becomes:

$$\frac{dN}{d\tau'} = ae^{-E} - \frac{cAN}{N+1} - bNe^{-E} \quad (4a)$$

$$\frac{dA}{d\tau'} = \frac{fNA}{N+1} - A \quad (4b)$$

$$\frac{dE}{d\tau'} = \epsilon(dA - E) \quad (4c)$$

135 We now let  $\epsilon \rightarrow 0$ . In the dimensional model, this is equivalent to  $\epsilon = \eta/\delta \rightarrow 0$ , which corresponds  
 136 to the assumption that algal loss occurs at a much faster rate than EPS decay ( $\delta \gg \eta$ ). In this limit,  
 137 the amount of EPS remains at its initial value,  $E(t) = E(0) = E_0$ . We now analyze the remaining  
 138 two equations, (4a) and (4b). Setting  $\frac{dN}{d\tau'} = \frac{dA}{d\tau'} = 0$ , we find two steady-state solutions:

$$p_0 := \left(\frac{a}{b}, 0\right)$$

$$p_1 := \left(\frac{1}{f-1}, \frac{f(af-a-b)}{e^{E_0}c(f-1)}\right).$$

139 Again, we find a trivial, algae-free solution  $p_0$ , as well as a possible nontrivial solution  $p_1$  with  
 140 persistent algae, provided  $\frac{a}{b} > \frac{1}{(f-1)} > 0$ , which is necessary to guarantee positive entries in  $p_1$ .  
 141 Evaluating the Jacobian of the matrix differential of (4a)-(4b) at  $p_0$  yields the following eigenvalues  
 142 (see S2):

$$\lambda_1 = -be^{-E_0}, \lambda_2 = \frac{af-a-b}{a+b}$$

143 Since  $\lambda_1$  will always be negative, the stability of  $p_0$  is determined by the sign of  $\lambda_2$ . The sign of  
 144  $\lambda_2$  hinges on the same inequality as the existence of  $p_1$ , specifically the relationship between  $\frac{a}{b}$  and  
 145  $\frac{1}{(f-1)}$ . If  $p_1$  exists, then  $p_0$  is unstable, since  $\frac{a}{b} > \frac{1}{(f-1)}$  and so  $\lambda_2 > 0$ . If  $p_1$  does not exist,  $p_0$  is  
 146 stable, since  $\frac{a}{b} < \frac{1}{(f-1)}$  and so  $\lambda_2 < 0$  (Figure 1).

147 Turning our attention to the stability of the nontrivial equilibrium, we evaluate the Jacobian at  
 148  $p_1$  (see S2):

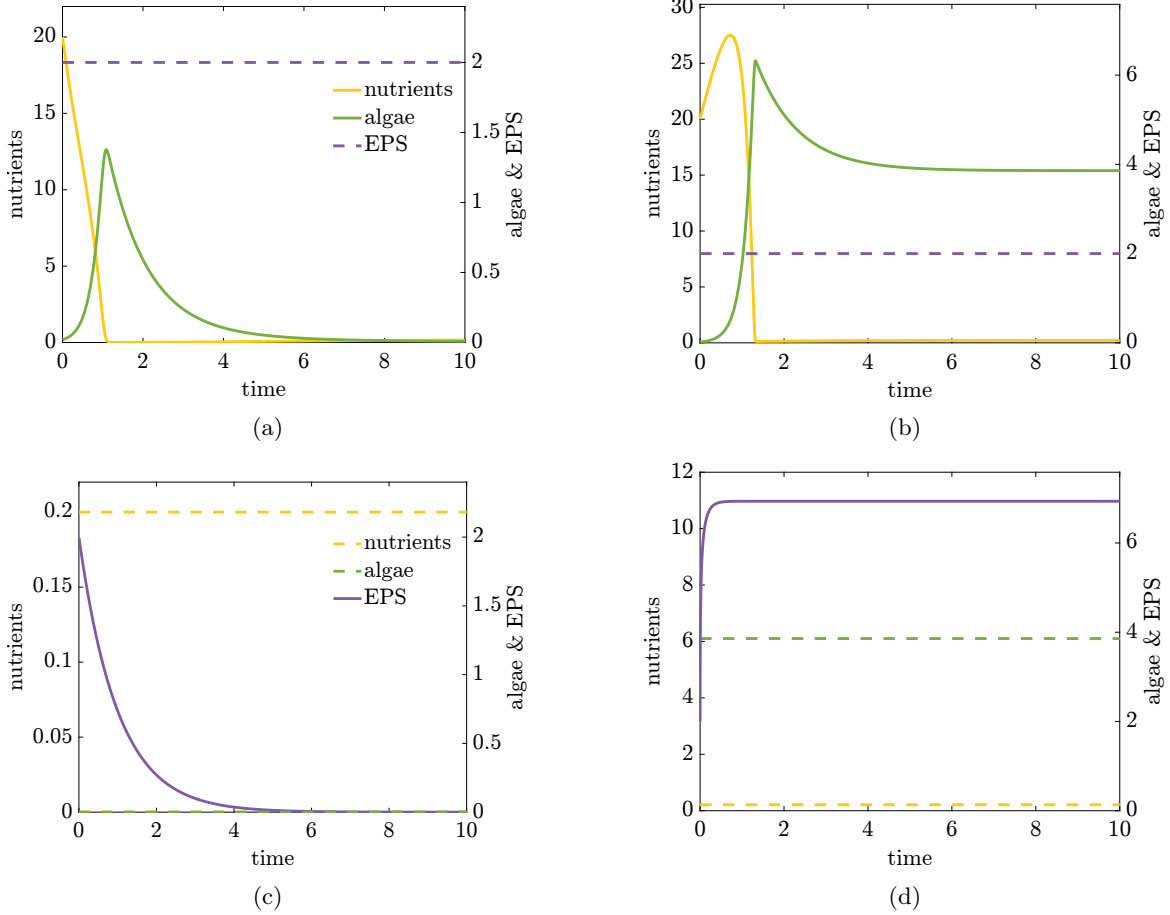
$$J_{p_1} = \begin{bmatrix} -\frac{e^{-E_0}(f^2a-2af+a+b)}{c} & -\frac{c}{f} \\ \frac{f}{c} & 0 \end{bmatrix}$$

149 We use the trace and determinant to determine stability. If  $\frac{a}{b} > \frac{1}{(f-1)}$ , then the trace of the  
 150 Jacobian is negative and the determinant is positive, confirming that  $p_1$  becomes a stable sink for  
 151 the same values of  $\frac{a}{b}$  at which  $p_0$  loses stability (Figure 1). This result echoes our findings in Section  
 152 2.1, where we identified a transcritical bifurcation governed by the balance between nutrient inflow  
 153 and outflow.

154 We also rule out the possibility of a Hopf bifurcation for this system. For a Hopf bifurcation, the  
 155 trace of the Jacobian would need to be zero:

$$\text{Trace}(J_{p_1}) = 0 \implies f^2a - 2af + a + b = 0 \implies \frac{a}{b} = \frac{-1}{(f-1)^2}$$

156 which is not possible for nonnegative values of  $a$  and  $b$ .



**Figure 1:** Two possible asymptotic dynamics arise under Regime 1, where we assume EPS dynamics are much slower than algal and nutrient dynamics. In both (a) and (b), nutrients and algae evolve on the fast timescale while EPS is held fixed at its quasi-steady state value, ( $E = E_0$ ). In both (a) and (b), nutrients and algae are assumed to instantaneously reach their fast quasi-steady state  $p_0$  and  $p_1$ , respectively, reducing the system to a single slow EPS equation. In both (a) and (c), parameters are such that nutrient loss exceeds nutrient input, resulting in algae and EPS asymptotically approaching 0. Dimensional parameters are as in Table 1, except for  $\phi = .0001$  and  $\psi = .05$ . Corresponding nondimensional parameters are  $a = 1.4, b = 7.1, f = 5.7$  thus  $\frac{a}{b} < \frac{1}{f-1}$ . In both (b) and (d), nutrient input is sufficient to sustain algae and EPS asymptotically. Parameters are as in Table 1, here nondimensional parameters  $a = 142.9, b = 1.4, f = 5.7$  thus  $\frac{a}{b} > \frac{1}{f-1}$ .

157 We now focus on the slow model dynamics of (3b). For this analysis, we do not rescale time, so  
 158 when we let  $\epsilon \rightarrow 0$ , the left-hand side of the equations in (3a) becomes zero. In this quasi-steady-  
 159 state limit, we find the same two nutrient and algal equilibria as in the fast system above,  $p_0$  and  
 160  $p_1$ . In the trivial quasi-steady state  $p_0$ , (3b) simplifies to  $\frac{dE}{d\tau} = -E$ , so the amount of EPS decays  
 161 exponentially from its initial condition  $E(0)$  in the absence of algae. This behavior is biologically  
 162 reasonable, as no significant amount of new EPS is produced in the absence of algae, and any existing  
 163 EPS decay. If we instead substitute the nontrivial quasi-steady state  $p_1$  into (3b), we have

$$\frac{dE}{d\tau} = \frac{fd(af - a - b)}{e^E c(f - 1)} - E. \quad (5)$$

164 This results in a nontrivial steady-state solution for EPS (Figure 1),

$$E^* = W\left(\frac{fd(af - a - b)}{c(f - 1)}\right)$$

165 where  $W(\cdot)$  is the Lambert W function. The Lambert W function is defined as the solution  $W(x)$  to  
 166 the transcendental equation  $We^W = x$  with real domain  $x \in [-\frac{1}{e}, \infty)$ . Note that in order for  $W(\cdot) > 0$   
 167 and an equilibrium point  $E^*$  to exist, the argument of the Lambert W function is nonnegative, which  
 168 requires  $af - a - b \geq 0$ , or equivalently  $\frac{a}{b} > \frac{1}{f-1}$ . To assess the local stability of  $E^*$ , we differentiate  
 169 the right-hand side of equation (5) with respect to  $E$  and evaluate at  $E^*$ , which yields:

$$\frac{d}{dE} \left( \frac{dE}{d\tau} \right) \Big|_{E=E^*} = -W \left( \frac{fd(af - a - b)}{c(f-1)} \right) - 1 \quad (6)$$

170 The stability of  $E^*$  is determined by the sign of (6), which depends on  $W \left( \frac{fd(af - a - b)}{c(f-1)} \right)$ . Since  
 171  $W(-\frac{1}{e}) = -1$ , a change in stability occurs when

$$\frac{fd(af - a - b)}{c(f-1)} = -\frac{1}{e}. \quad (7)$$

172 We can reframe this relationship in terms of the ratio of nutrient fluxes,  $\frac{a}{b}$ ,

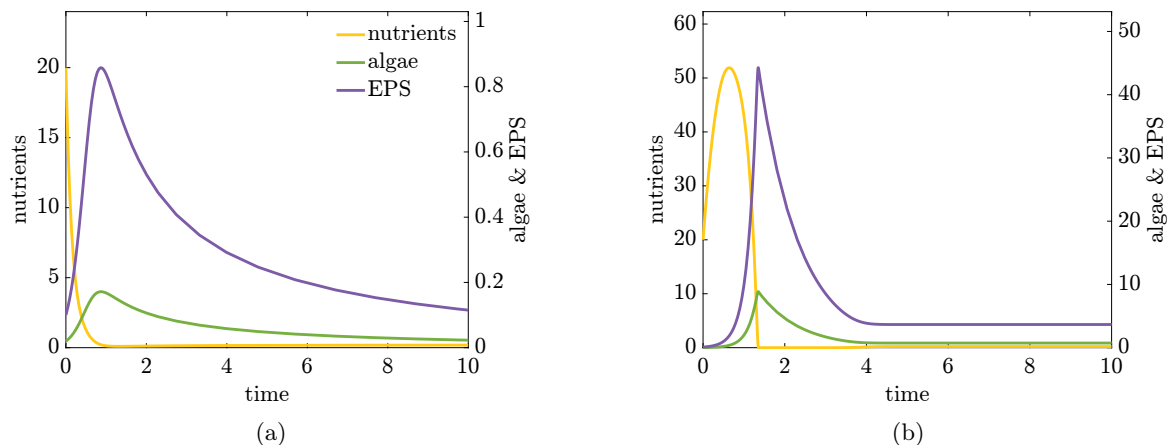
$$\frac{a}{b} = \frac{-c}{efdb} + \frac{1}{f-1}$$

173 When  $\frac{a}{b}$  is greater than this threshold, then  $E^*$  in the slow model is stable (Figure 1).

174 In the fast model, algae persist when  $\frac{a}{b} > \frac{1}{(f-1)}$ . Because (5) assumes the same nontrivial algal  
 175 steady state as in the fast model, if

$$\frac{a}{b} > \frac{1}{f-1} > \frac{-c}{efbd} + \frac{1}{f-1},$$

176 then, algae in the fast model will rapidly reach a nontrivial steady state, with a corresponding  
 177 nontrivial EPS steady state in the slow model. This fast-slow analysis reveals dynamics consistent  
 178 with our previous findings: if nutrient input rates are sufficiently high relative to nutrient loss, the  
 system can sustain positive algae and EPS concentrations asymptotically.



**Figure 2:** Two possible asymptotic regimes under Regime 2, arising from timescale separation in which EPS dynamics track algal dynamics. In this regime, EPS adjusts on the same fast timescale as algal and nutrient dynamics. In both (a) and (b), parameter values transition from a regime in which nutrient loss exceeds nutrient input ( $\frac{a}{b} < \frac{1}{f-1}$ ), resulting in algal extinction and stability of the trivial equilibrium  $p_0$ , to a regime in which nutrient input dominates loss ( $\frac{a}{b} > \frac{1}{f-1}$ ), yielding a stable positive algal equilibrium  $p_1$ . (a): Dimensional parameters are as in Table 1, except for  $\phi = .0001$  and  $\psi = .05$ . for the left column. Here, nondimensional parameters are  $a = 1.4, b = 7.1, f = 5.7$  thus  $\frac{a}{b} < \frac{1}{f-1}$ . (b): Parameters are as in Table 1, here, nondimensional parameters are  $a = 142.9, b = 1.4, f = 5.7$  thus  $\frac{a}{b} > \frac{1}{f-1}$ .

180 **2.2.2 Regime 2: EPS dynamics are fast, responding instantaneously to**  
 181 **changes in algal biomass**

182 In the previous section, we considered a possible regime where EPS dynamics are much slower  
 183 than those of nutrients and algae. Here, we consider the implications of the opposite extreme: EPS  
 184 responds instantaneously to changes in nutrient and algae levels. Mathematically, we model this  
 185 by assuming EPS perfectly tracks algae, so  $E(t) = \sigma A(t)$ , where  $\sigma$  is a scaling term. Under this  
 186 assumption, model (2) becomes

$$\begin{aligned}\frac{dN}{dt} &= \phi e^{-\frac{\sigma A}{\mu}} - \frac{\nu N}{\gamma + N} A - \psi e^{-\frac{\sigma A}{\mu}} N \\ \frac{dA}{dt} &= \xi \frac{\nu AN}{\gamma + N} - \delta A \\ E(t) &= \sigma A(t).\end{aligned}\tag{8}$$

187 Nondimensionalizing the system (Appendix 8) yields:

$$\begin{aligned}\frac{dN}{dt} &= a e^{-hA} - \frac{cNA}{N+1} - bN e^{-hA} \\ \frac{dA}{dt} &= \frac{fNA}{N+1} - A \\ E(t) &= hA\end{aligned}\tag{9}$$

188 We ignore  $E(t)$  in our analysis, since  $E(t)$  dynamics are directly proportional to  $A(t)$ . This system  
 189 again has two steady-state solutions for  $N$  and  $A$ :

$$\begin{aligned}r_0 &= \left(\frac{a}{b}, 0\right) \\ r_1 &= \left(\frac{1}{f-1}, \frac{W\left(hf \frac{af-a-b}{c(f-1)}\right)}{h}\right),\end{aligned}$$

190 corresponding to a trivial, algae-free steady state  $r_0$ , and a nontrivial steady state  $r_1$  with persistent  
 191 algae (and thus persistent EPS at concentration  $E^* = hA^*$ ).

192 Evaluating the Jacobian of the matrix differential equation of (9) at  $r_0$  yields the following  
 193 eigenvalues (see S3):

$$\lambda_1 = -b, \quad \lambda_2 = \frac{af - a - b}{a + b}.$$

194  $\lambda_1$  is always negative, so the stability of  $r_0$  is determined by the sign of  $\lambda_2$ . As in Section 2.2.1,  
 195 we find that  $r_0$  loses stability through a transcritical bifurcation at  $\frac{a}{b} = \frac{1}{(f-1)}$ , demonstrating the  
 196 importance of nutrient fluxes in sustaining persistent algal communities.

197 Observe that this transcritical bifurcation occurs exactly at the point where the entries of  $r_1$   
 198 become nonnegative.

199 Due to the presence of the Lambert W function, analyzing the stability of  $r_1$  is not as straight-  
 200 forward as evaluating the Jacobian of (9) at  $r_1$ . Instead, we proceed as follows. First, we substitute  
 201  $N^* = \frac{1}{f-1}$  into (9) to obtain an expression for  $A^*$  as a function of system parameters:

$$\frac{dN}{dt} = a e^{-hA^*} - \frac{cA^*}{f} - \frac{b e^{-hA^*}}{f-1} = 0.$$

202 We can then rearrange this expression in terms of one of our parameters, here chosen to be  $b$ , the  
 203 physical nutrient loss rate,

$$b = (f-1)(a - cA^* e^{hA^*} / f).\tag{10}$$

204 Plugging this expression into both  $\frac{dN}{dt}$  and  $\frac{dA}{dt}$  in (9), we can evaluate the Jacobian of (9), finding  
 205 the determinant and trace as functions of  $A^*$  (see S3):

$$\text{Det}(J_{r_1}) = \frac{(hA^* + 1)cA^*(f-1)^2}{f^2}\tag{11}$$

$$Tr(J_{r_1}) = \frac{(f-1)(-e^{-hA^*}af^2 + cA^*)}{f^2} \quad (12)$$

206 The determinant is always nonnegative, and it can be proven that the trace is always negative  
 207 (Appendix 8), so this equilibrium is stable.

Next, we ask whether a Hopf bifurcation could occur by examining the condition  $Tr(J_{r_1}) = 0$ . Setting the trace to zero in (12) gives:

$$cA^* = af^2e^{-ha}$$

208 Substituting this term into the expression for  $b$  in equation (10), we compute:

$$\begin{aligned} \text{sign}(b) &= \text{sign}(-cA^* + afe^{-hA^*}) \\ &= \text{sign}(-af^2e^{-hA^*} + afe^{-hA^*}) \\ &= \text{sign}(-afe^{-hA^*}(f+1)) \\ &< 0 \end{aligned}$$

209 Again, we see that achieving  $Tr(J_{r_1}) = 0$  requires  $b < 0$ , which is biologically implausible. Hence, for  
 210 all positive parameter values,  $Tr(J_{r_1}) < 0$ , and the equilibrium point  $r_1$  is a stable sink (Figure 2).

### 211 2.2.3 Regime 3: EPS dynamics occur over similar timescales to those of algae 212 and nutrients

213 The final regime explores the hypothesis that EPS dynamics occur on a timescale similar to that of  
 214 algae and nutrients, requiring us to consider the full system of three ordinary differential equations  
 215 in (2). We have already nondimensionalized this model in (3), and we now analyze the full system  
 216 dynamics rather than making a quasi-steady-state assumption, as was done in Section 2.1. The  
 217 steady states  $(N^*, A^*, E^*)$  of (3) are:

$$\begin{aligned} q_0 &= \left(\frac{a}{b}, 0, 0\right) \\ q_1 &= \left(\frac{1}{f-1}, \frac{1}{d}W\left(\frac{(af-a-b)fd}{c(f-1)}\right), W\left(\frac{(af-a-b)fd}{c(f-1)}\right)\right) \end{aligned}$$

218 Notice the similarity to the equilibria  $p_0$  and  $p_1$  in Regime 1 (Section 2.2.1) and  $r_0$  and  $r_1$  in Regime  
 219 2 (Section 2.2.2). As in Section 2.2.2, the nontrivial EPS steady state is a scalar multiple of the algal  
 220 steady state. Note that  $N^*$  in  $q_1$  is only plausible if  $f \geq 1$ . As above in Regime 1 (Section 2.2.1),  
 221 this transcritical bifurcation occurs exactly at the point where the Lambert W entries of  $q_1$  become  
 222 nonnegative.

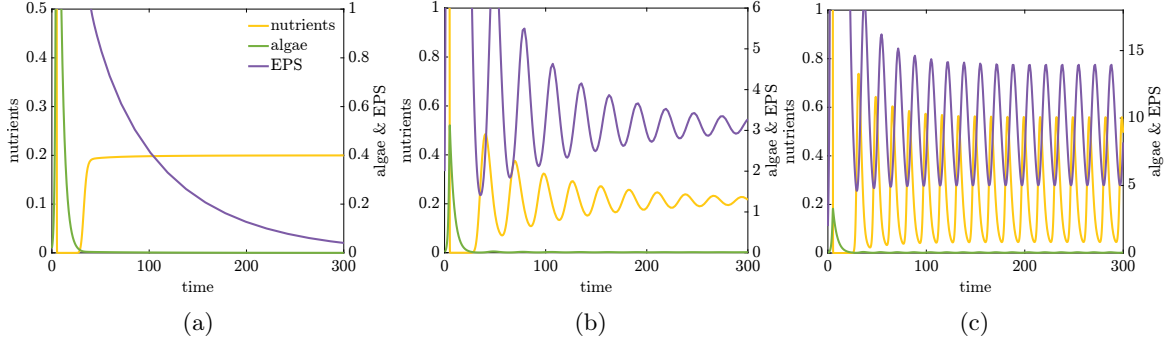
To determine the stability of  $q_0$ , we examine the eigenvalues of the Jacobian of (3) at  $q_0$ , (see S4):

$$\lambda_1 = -1, \lambda_2 = -b, \lambda_3 = \frac{af-a-b}{a+b}.$$

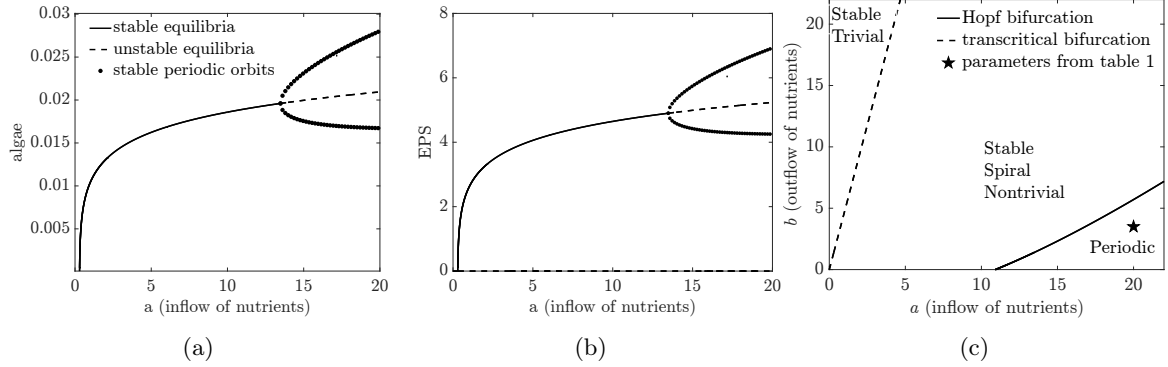
223 As in the previous two regimes, the trivial steady state  $q_0$  is stable if  $\frac{a}{b} < \frac{1}{(f-1)}$ , and it is the ratio  
 224 of nutrient input to loss that determines algal persistence (Figure 3).

225 We next examine the stability of  $q_1$ , where algae and EPS persist. Numerical bifurcation diagrams  
 226 suggest that immediately following the transcritical bifurcation, the nontrivial steady state  $q_1$  is a  
 227 stable spiral, and that increasing nutrient input far enough leads not only to persistent algae and EPS  
 228 but may eventually result in periodic dynamics through a Hopf bifurcation (Figure 3). Motivated  
 229 by these numerical findings, we look for two bifurcations: one where the eigenvalues of the Jacobian  
 230 evaluated at the nontrivial fixed point become negative, signifying the stability of  $q_1$ ; and one where  
 231 a complex conjugate pair of eigenvalues becomes purely imaginary, signifying periodic asymptotic  
 232 dynamics.

233 In sections 2.2.1 and 2.2.2, the system was two-dimensional, allowing stability to be determined  
 234 from the trace and determinant of the Jacobian. In the full three-dimensional model, this approach is  
 235 no longer sufficient. Moreover,  $q_1$  involves the Lambert W function, which makes direct substitution



**Figure 3:** Three possible asymptotic behaviors under Regime 3, arising from the ratio of nutrient inflow to nutrient outflow. **(a):** When the ratio of nutrient input to loss is too low ( $\frac{a}{b} < \frac{1}{f-1}$ ), then the system is asymptotically free from algae and EPS ( $q_0$  is stable). Parameters are as in Table 1, except  $\phi = .0001$  and  $\psi = .05$  which corresponds to nondimensional parameters  $a = 1.4$ ,  $b = 71.4$ ,  $f = 5.7$ . **(b):** As  $\frac{a}{b}$  increases and crosses the transcritical bifurcation threshold, we get a stable spiral towards the nontrivial equilibrium,  $q_1$ . Note that algae are also oscillatory here, but difficult to see due to the low levels at these parameter values. Parameters are as in Table 1, except  $\phi = .0001$  and  $\psi = .001$  with nondimensional parameters  $a = 1.4$ ,  $b = .14$ ,  $f = 5.7$ . **(c):** When  $(a, b)$  is on the right side of the manifold (Figure 4), we get a limit cycle around equilibrium point  $q_1$ . As in the middle figure, algae are oscillatory here, though difficult to see. Here, parameters are as in Table 1.



**Figure 4:** Bifurcation diagrams for regime 3. **(a):** Bifurcation diagram of asymptotic algal biomass,  $A$ . **(b):** Bifurcation diagram of asymptotic EPS concentration,  $E$ . **(c):** Bifurcation diagram of both inflow of nutrients,  $a$ , and outflow of nutrients,  $b$ . For values of  $a$  slightly larger than 0, a transcritical bifurcation occurs in which the trivial equilibrium loses stability and a stable nonzero equilibrium emerges. As  $a$  increases further, the nontrivial equilibrium undergoes a Hopf bifurcation, giving rise to stable periodic oscillations. Parameters are as in Table 1.

236 into the Jacobian algebraically cumbersome. For this reason, we instead work symbolically with the  
 237 Jacobian and impose steady-state relationships later in the calculation, thereby avoiding explicit  
 238 Lambert W terms.

239 To explore the possibility of a periodic solution, we construct the full Jacobian  $J_{q_1}$ , evaluate it  
 240 at the fixed point  $q_1$  (see S4), and analyze its eigenvalues. We construct a modified characteristic  
 241 polynomial,  $p_{char} = \det(J_{q_1} - i\sqrt{\omega}I)$ , where  $i\sqrt{\omega}$  are the purely complex eigenvalues of the system.  
 242 The roots of the characteristic polynomial may contain both real and imaginary parts. At a Hopf  
 243 bifurcation, we require both parts to vanish simultaneously. Therefore, we set  $R = \text{Re}(p_{char}) = 0$   
 244 and  $\text{Im}(p_{char}) = 0$ . We first solve  $\text{Im}(p_{char}) = 0$ , which gives an expression for  $\omega$  (see S4). Next,  
 245 we consider  $\text{Re}(p_{char}) = 0$ , but solving this equation directly for  $\omega$  is challenging because the state  
 246 variables  $N$ ,  $A$ , and  $E$  appear in the expression. To address this challenge, we impose the steady-state  
 247 conditions from the nondimensional system (3), so

$$E^* = dA^*$$

248 and

$$N^* = \frac{1}{f-1}.$$

249 Substituting these terms into the differential equation for  $\frac{dN}{dt}$  yields an expression that depends on  
 250 several parameters and the state variable  $A$ :

$$\frac{dN}{dt} = \frac{f(af - a - b)e^{-dA^*} - cA^*(f-1)}{f(f-1)}.$$

251 Setting  $\frac{dN}{dt} = 0$  gives the steady state relationship between parameters and  $A$ ,

$$b = -\frac{(f-1)(cA^*e^{dA^*} - af)}{f}.$$

252 Now, with both  $\omega$  and  $b$  expressed analytically (See S4), we can find the solutions to  $Re(p_{char}) = 0$ .  
 253 We choose to express these solutions in terms of  $a$ , maintaining our interest in how nutrient fluxes  
 254 shape system dynamics. The parameter  $a$  appears quadratically in  $Re(p_{char})$ , and we find a pair of  
 255 solutions: one positive and one negative. The negative root is irrelevant, so the remaining solution  
 256 gives an analytical expression for  $a$  as a function of several parameters and  $A^*$  (see S4). Together,  
 257 these expressions for  $a$  and  $b$  describe a Hopf bifurcation manifold that determines whether the  
 258 nontrivial steady state is stable or whether there is a stable periodic orbit around it (Figure 4).

### 259 2.3 Comparison between biological regimes

Regime	Trivial steady state ( $N^*, A^*, E^*$ )	Stability Condition	Nontrivial steady state ( $N^*, A^*, E^*$ )	Hopf Bifurcation
Regime 1	$(\frac{a}{b}, 0, 0)$	$\frac{a}{b} < \frac{1}{f-1}, (f > 1)$	$(\frac{1}{f-1}, \frac{f(af-a-b)}{e^{E_0}c(f-1)}, W(\frac{d^f(af-a-b)}{c(f-1)}))$	N/A
Regime 2	$(\frac{a}{b}, 0, 0)$	$\frac{a}{b} < \frac{1}{f-1}, (f > 1)$	$(\frac{1}{f-1}, \frac{1}{h}W(h\frac{f(af-a-b)}{c(f-1)}), hA^*)$	N/A
Regime 3	$(\frac{a}{b}, 0, 0)$	$\frac{a}{b} < \frac{1}{f-1}, (f > 1)$	$(\frac{1}{f-1}, \frac{1}{d}W(d\frac{f(af-a-b)}{c(f-1)}), dA^*)$	Figure 4

**Table 2:** Summary comparing asymptotic model behavior across regimes.

260 Across all three regimes, the trivial steady state, without persistent algae or EPS, is stable when  
 261 the ratio of nutrient inflow to nutrient loss is too small (i.e.  $\frac{a}{b} < \frac{1}{f-1}$  for  $f > 1$ ) (Table 2). In this  
 262 case, nutrient inflow  $a$  is too low to offset physical and biological losses (captured in  $b$ , the nutrient  
 263 loss rate, and  $f$ , the algal nutrient uptake rate), preventing algal establishment and resulting in EPS  
 264 production. Conversely, when nutrient inputs are sufficiently high relative to loss ( $\frac{a}{b} > \frac{1}{f-1}$ ), algae  
 265 persist, with resulting sustained EPS production.

266 The parameters  $h$  and  $d$  seen in the nontrivial steady state in Regimes 2 and 3 (Table 2) share the  
 267 same units, and both represent the growth rate of EPS. In Regime 2, the parameter  $\sigma$  represents the  
 268 efficiency with which algal biomass is converted into EPS, setting the proportionality between algal  
 269 growth and EPS accumulation. Larger values of  $\sigma$  correspond to more rapid EPS production per  
 270 unit of algal biomass, while smaller values indicate weaker coupling between algal growth and EPS  
 271 synthesis. When  $\sigma$  is set to  $\sigma = \frac{\rho}{\eta}$ , where  $\rho$  is the production rate of EPS and  $\eta$  is the degradation  
 272 of EPS, then  $d = h$ , leading to identical steady-state behavior in Regimes 2 and 3. In both of these  
 273 regimes, EPS concentration is directly proportional to algal biomass at steady state.

274 Based on the parameters found in the literature (Hunke et al. (2024); Krembs et al. (2011)),  
 275 we have  $E^* = dA^* > A^*$ , indicating that  $d, h > 1$ . Therefore, in both Regimes 2 and 3, the  
 276 nontrivial equilibrium EPS concentration exceeds the corresponding algal biomass, consistent with  
 277 observational studies (Stoecker et al. (2000)).

278 To compare the nontrivial algal steady states across the different regimes in Table 2, we begin  
 279 with the expression

$$x = \frac{f(af - a - b)}{c(f - 1)} > 0$$

280 which appears in each regime’s nontrivial steady state  $A^*$ . First, we compare the value of Regime 1,  
 281  $xe^{-E_0}$ , to that of Regime 2,  $\frac{1}{h}W(hx)$ . To determine which is larger, we find when  $\frac{1}{h}W(hx) = xe^{-E_0}$ .  
 282 Solving for  $x$  yields  $x = \frac{E_0 e^{E_0}}{h}$ .

283 When  $0 < x < \frac{E_0 e^{E_0}}{d}$ , the nontrivial steady-state algal biomass in Regime 2 exceeds that in  
 284 Regime 1, i.e.,  $A_2^* > A_1^*$ . At the threshold value  $x = \frac{E_0 e^{E_0}}{d}$ , the two regimes yield identical steady-  
 285 state algal biomass, so that  $A_1^* = A_2^*$ . For  $x > \frac{E_0 e^{E_0}}{d}$ , the ordering reverses, and Regime 1 supports  
 286 a larger asymptotic algal biomass than Regime 2, i.e.,  $A_1^* > A_2^*$ . An identical comparison holds for  
 287 Regime 3, as its expression for  $A^*$  differs from that of Regime 2 only by replacing  $h$  with  $d$ , leading  
 288 to the same conclusions regarding the relative magnitudes of the nontrivial steady states. Finally, if  
 289  $E_0 > 0$ , the quantity  $\frac{x}{e^{E_0}} < x$  indicates that regime 1 yields less asymptotic algal biomass than the  
 290 unmodified baseline value  $x$ ; in this case, EPS reduces the nontrivial steady state algal concentration  
 291  $A^*$ .

292 In summary, the comparison across regimes reveals a clear hierarchy governed by a threshold  
 293  $x = \frac{E_0 e^{E_0}}{d}$ : Below this threshold, Regimes 2 and 3 support higher asymptotic algal biomass than  
 294 Regime 1, while above it, the ordering reverses. Regime 2 serves as an intermediate case, aligning  
 295 with Regime 3 when  $d = h$  and shifting otherwise. In all cases, EPS reduces steady-state algal  
 296 biomass relative to the baseline value  $x$ , with the magnitude of this reduction depending on EPS  
 297 production parameters.

## 298 2.4 Transient bloom dynamics

299 In addition to studying the asymptotic behavior of algae and EPS, we are also interested in their  
 300 transient dynamics during an algal bloom. Transient dynamics occur over shorter, often ecologically  
 301 important time scales. Our transient analysis is inspired by that of Huppert (2002), who explored  
 302 transient bloom dynamics for a classic model of nutrient-driven phytoplankton blooms (Huppert  
 303 et al. (2002)). The relative simplicity of that model allowed for an analytic analysis of the transient  
 304 dynamics (Huppert et al. (2002)). Adding EPS to the nutrient-algae system precludes the use of those  
 305 analytic approaches, so we instead numerically investigate the dependence of the bloom dynamics on  
 306 initial conditions, nutrient fluxes, and EPS accumulation and decay rates. To quantify these effects,  
 307 we conducted a local sensitivity analysis to determine the parametric dependencies of the maximum  
 308 nutrient, algal, and EPS concentrations ( $N_{max}$ ,  $A_{max}$ , and  $E_{max}$ ) during the algal bloom. We focus  
 309 our transient analysis on the model under Regime 3, which makes no strong assumptions about the  
 310 timescale of EPS dynamics.

311 Note that we define a “bloom” as occurring whenever there is an initial increase in algae imme-  
 312 diately following time  $t = 0$ . Mathematically, this requires  $\frac{dA}{dt}|_{t=0} > 0$ , which necessitates  $N_0 > \frac{1}{f-1}$   
 313 in (3). In what follows, we assume  $N_0$  is above this threshold, resulting in an initial increase in  $A(t)$ .  
 314 We then vary parameters by 50-150% in the dimensional model (2) around the values in Table 1.

### 315 *Influence of nutrient fluxes on bloom peaks*

316 Previous work on nutrient-driven phytoplankton dynamics suggests that an increase in the nutrient  
 317 inflow rate leads to higher phytoplankton and nutrient peaks (Huppert et al. (2002)), and vice versa  
 318 for the nutrient loss rate. Consistent with these findings, we observe that increasing  $\phi$  results in higher  
 319 maximum concentrations of algae ( $A_{max}$ ), nutrients ( $N_{max}$ ), and EPS ( $E_{max}$ ), while increasing  $\psi$   
 320 lowers them (Figure 5).

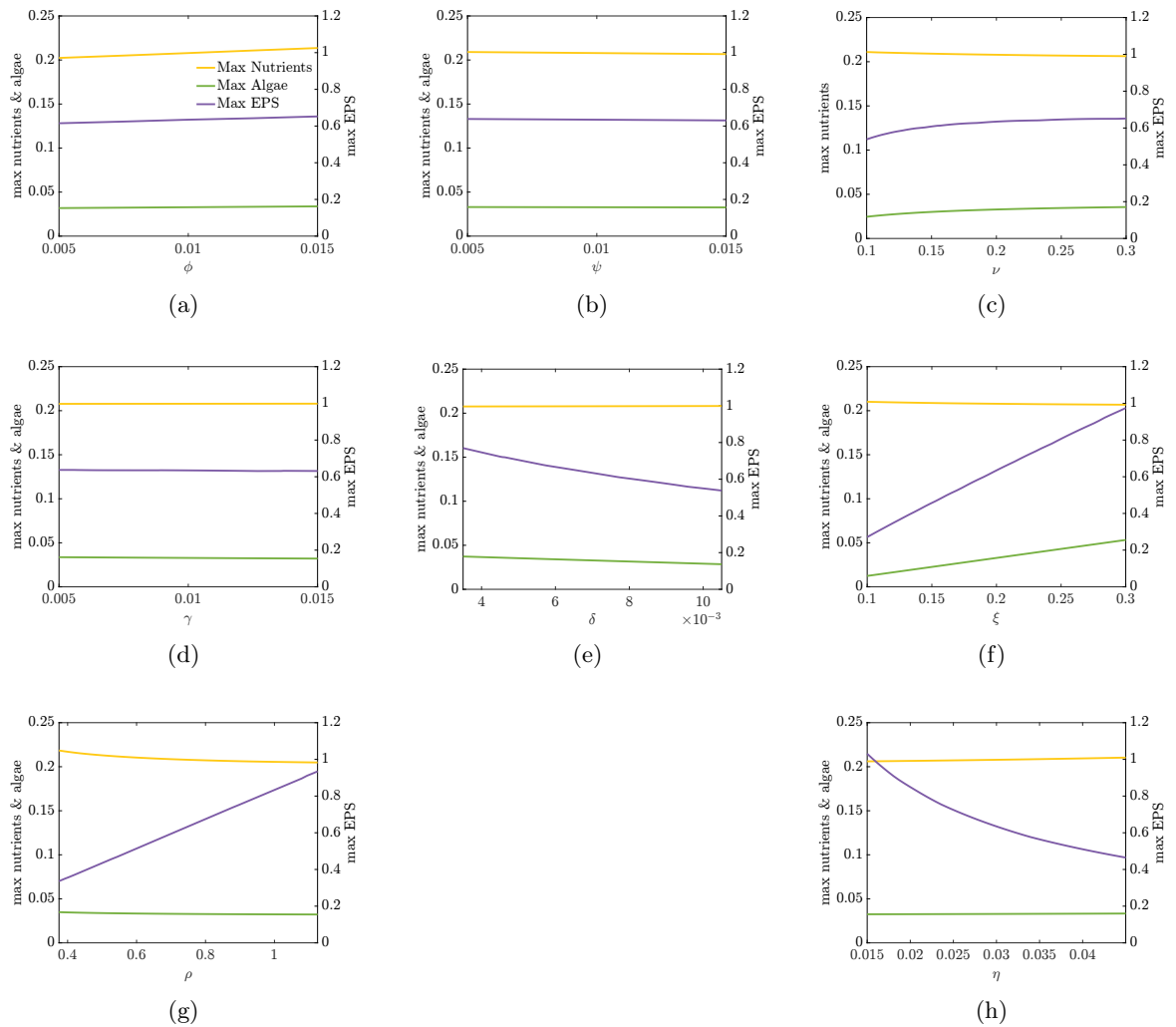
### 321 *Influence of algal growth and loss rates on bloom peaks*

322 As the nutrient uptake rate of bottom-ice algae,  $\nu$ , increases,  $N_{max}$  decreases, while  $A_{max}$  and  $E_{max}$   
 323 both increase, as one may have intuitively expected (Figure 5). Similarly, decreasing the nitrogen  
 324 half-saturation constant,  $\gamma$ , should increase the rate of nutrient uptake by algae, but within the tested  
 325 range  $\gamma$  has little effect on bloom peaks (Figure 5). In contrast, increasing the algal mortality rate,  $\delta$ ,  
 326 produces the expected effect: both  $A_{max}$  and  $E_{max}$  decline, while  $N_{max}$  remains largely unchanged

327 (Figure 5). Finally, increasing the chl-a to nitrogen conversion ratio,  $\xi$ , increases both algal and EPS  
 328 peaks, accompanied by a slight reduction in the nutrient maximum (Figure 5).

### 329 *Influence of EPS accumulation and loss rates on bloom peaks*

330 EPS influences bloom dynamics in ways consistent with the hypothesized EPS-mediated feedback  
 331 mechanism. Increasing the EPS production rate  $\rho$  leads to a higher maximum EPS concentration  
 332 ( $E_{max}$ ), while reducing the peak of the algal bloom ( $A_{max}$ ) (Figure 5). Similarly, increasing the  
 333 degradation rate of EPS  $\eta$  causes  $E_{max}$  to decrease while  $N_{max}$  and  $A_{max}$  increase (Figure 5). These  
 334 inverse relationships between peak EPS accumulation and peak bloom intensity suggest that the  
 335 negative EPS-mediated biophysical feedback, which is the focus of this paper, reduces peak algal  
 336 concentrations, in addition to the reduction in asymptotic algal concentrations explored earlier.

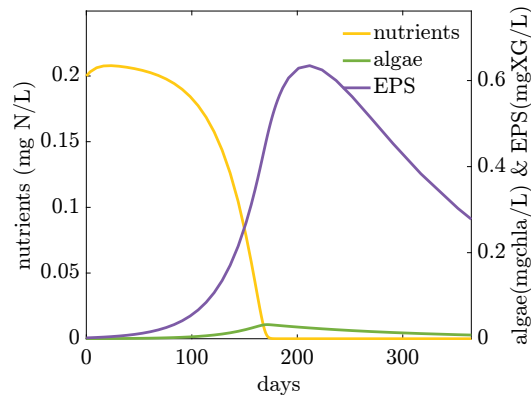


**Figure 5:** Each graph represents the effects that each parameter has on the nutrient, algal, and EPS concentration peaks ( $N_{max}$ ,  $A_{max}$ ,  $E_{max}$  respectively). Parameter values vary between 50% and 150% of the values in Table 1. In contrast to the other figures, the algal peak  $A_{max}$  is plotted on the left y-axis (rather than the right y-axis) to show its variation more clearly.

## 337 3 Real world model implications

338 We selected parameter values consistent with observations in polar sea ice systems (Table 1). These  
 339 values include both empirically measured quantities and reasonable estimates derived from biogeo-  
 340 chemical literature, and are intended to represent plausible conditions encountered within sea ice.

341 Figure 6 shows the solution of model 2 using the parameter values in Table 1. Here we see that high  
 342 nutrient concentrations initially support rapid algal growth and subsequent EPS production. Even-  
 343 tually, as nutrients are depleted, the bloom ends and the system transitions to a nutrient-limited  
 state (Collins et al. (2008); Krembs et al. (2002)).



**Figure 6:** Time series of nutrients, algae, and EPS over the first simulated year. Algal biomass and EPS persist while nutrient concentrations are gradually depleted to a sustained level, consistent with observations found from Arctic sea ice ecosystems. The modeled concentrations of nutrients, algae, and EPS during a bloom fall within reported ranges in the literature, indicating a realistic parameterization. However, the duration of the simulated algal bloom happens later than typically observed in field studies. Parameter values are as in Table 1.

344

## 345 4 Discussion

346 In this paper, we extended a classic nutrient-algae model to include EPS-mediated biophysical feed-  
 347 backs. We considered three regimes to distinguish between possible timescales of EPS dynamics. All  
 348 regimes share two asymptotic scenarios: one in which EPS, algae, and nutrients persist, and another  
 349 in which, despite sustained nutrient availability, algae and EPS do not persist. Notably, the shift  
 350 between these two asymptotic states depends only on a combination of the nutrient fluxes and the  
 351 growth rate of algae; the bifurcation is unrelated to EPS dynamics, and is consistent with the bi-  
 352 furcation found in the nutrient-phytoplankton model, without EPS (Huppert et al. (2002)). In all  
 353 regimes, reductions in permeability caused by EPS accumulation lead to lower asymptotic nutrient  
 354 availability, ultimately suppressing algal growth and altering system stability. When the feedback is  
 355 sufficiently strong, this results in periodic oscillations in nutrients, algae, and EPS, highlighting the  
 356 role of EPS in determining whether the system approaches a steady state or oscillatory behavior.  
 357 Thus, across regimes, the system’s qualitative behavior remains consistent, even as the quantitative  
 358 solutions differ slightly. This robustness suggests that the essential feedback mechanics governing  
 359 EPS and algal interactions are preserved regardless of the timescale assumptions on EPS. Analysis  
 360 of model transient dynamics revealed that EPS-mediated biophysical feedbacks also reduce bloom  
 361 intensity, suggesting that these feedbacks have both short and long-term implications for primary  
 362 productivity within the ice.

363 Perhaps the most salient unknown in our models is which functional form best describes the effects  
 364 of EPS on nutrient fluxes. In addition to the exponential form analyzed here, we also considered  
 365 an alternate, Hill function for the nutrient flux terms ( $\alpha(E)$  and  $\beta(E)$ ) (Appendix 8). This analysis  
 366 demonstrates that varying the feedback strength encoded by the Hill function can qualitatively  
 367 alter the system’s dynamics, including the emergence of a Hopf bifurcation that is absent under  
 368 weaker feedback. Future empirical and modeling studies are needed to improve our confidence in the  
 369 functional forms for  $\alpha(E)$  and  $\beta(E)$ .

370 While our asymptotic analyses suggest EPS suppresses asymptotic algal concentrations and may  
 371 even cause periodic solutions, real sea ice ecosystems are governed by significant additional com-  
 372 plexities such as seasonal temperature fluctuations. These temperature swings shape ice and snow  
 373 dynamics, which determine the amount of light available for photosynthesis within the ice and the  
 374 availability and stability of the ice substrate as an algal habitat (Riedel et al. (2006)). Additional

375 biological interactions, including those with bacteria, heterotrophic grazers, and viruses, as well as  
376 competition with other organisms for light and nutrients, introduce further complexity not consid-  
377 ered here. Future modeling efforts could merge models of these more complex dynamics with the  
378 EPS-mediated feedbacks explored here.

379 This modeling study represents a novel approach to integrating dynamic sea-ice biology and  
380 physics at the microscale. Previous studies have treated these processes independently, focusing on  
381 either sea ice biogeochemistry or physics within sea ice (Lannuzel et al. (2020); Miller et al. (2015);  
382 Roukaerts et al. (2021); Tedesco and Vichi (2014)). This study is, to the best of our knowledge,  
383 the first to quantitatively investigate the influence of EPS on sea ice permeability and the resulting  
384 impacts on algal growth. The EPS-mediated feedbacks explored here suppressed algal biomass both  
385 during the bloom and asymptotically, while simultaneously reducing ice permeability, highlighting  
386 the tight coupling between sea ice permeability and biological productivity.

## 387 5 Supporting information

388 All supplemental information, analysis, and code used to generate all figures are available in a [GitHub](#)  
389 [repository](#). A corresponding DOI will be created following acceptance for publication.

## 6 Author contributions

ARJ led all aspects of the study, including model development, analysis, and manuscript preparation. JRR contributed substantially to the study's conceptualization and to the initial manuscript draft. All other authors contributed to model formulation, manuscript editing and revisions. ARJ acknowledges the use of an AI-based language tool solely for editing and polishing the manuscript text (spelling, grammar, and general style).

## 7 Acknowledgments

ARJ gratefully acknowledges support from the University of Utah Department of Mathematics and the U.S. National Science Foundation through Grant DMS-2136198. KMG gratefully acknowledges support from the US National Science Foundation through grants DMS-2136198 and DMS-2206171, and the Applied and Computational Analysis Program at the US Office of Naval Research through grants N00014-21-1-2909 and N00014-26-1-2114.

## 8 Declerations

**Competing Interests** The authors declare that they have no financial or non-financial interests that are directly or indirectly related to the work submitted for publication.

## Appendix A: Hill functional form

Based on our findings above that EPS-mediated feedback can result in periodic solutions, we now ask what level of negative feedback is required for the system to exhibit this periodic behavior. To explore the impacts of an alternative functional choice for the nutrient flux terms, we introduce an alternative functional form for  $\alpha(E)$  and  $\beta(E)$ . The Hill function captures saturation effects by modeling how increasing EPS diminishes feedback strength on nutrient inflow and outflow. This reflects some intuition that, beyond a certain concentration, additional EPS may have little additional impact on nutrient dynamics. We set  $\alpha(E) = \frac{\phi}{(1+\mu E)^p}$  and  $\beta(E) = \frac{\psi}{(1+\mu E)^p}$  in (1), where  $p$  controls the level of negative feedback. Nondimensionalization yields

$$\epsilon \frac{dN}{dt} = \frac{a}{(1+E)^p} - \frac{cNA}{1+N} - \frac{bN}{(1+E)^p} \quad (13a)$$

$$\epsilon \frac{dA}{dt} = \frac{fNA}{1+N} - A \quad (13b)$$

$$\frac{dE}{dt} = dA - E \quad (13c)$$

Note that the system is nondimensionalized with the same parameters as before (see S6) to aid in interpretation. This system, like the previous ones, has a trivial and nontrivial equilibrium,

$$g_0 = \left( \frac{a}{b}, 0, 0 \right)$$

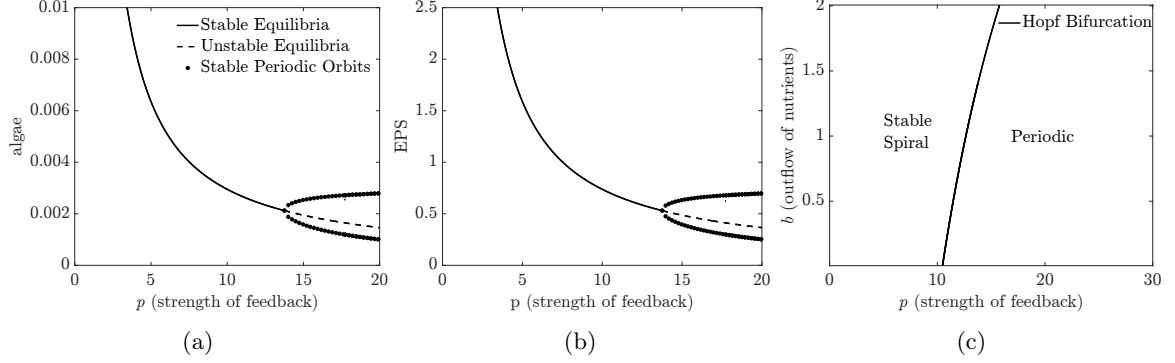
$$g_1 = \left( \frac{1}{f-1}, \frac{\left( \frac{c}{f-1} - a \right) \left( 1 + \frac{1}{f-1} \right)}{\left( E+1 \right)^p \left( \frac{c}{f-1} \right)}, Ad \right)$$

We evaluate the Jacobian of (13) at  $g_0$  and find the eigenvalues:

$$\mu_1 = -1, \mu_2 = -b, \mu_3 = \frac{fa - a - b}{b + a}.$$

As before, the stability of the trivial solution  $g_0$  is lost when  $\frac{a}{b} > \frac{1}{f-1}$ . The remaining stability analysis continues to follow the analysis of Section 2.2.3 (see S6). We highlight here how the results of model (13) differ from those of model (2), which had exponential terms for  $\alpha(E)$  and  $\beta(E)$ . In this model, we can vary  $p$ , the strength of the EPS-mediated feedback in the system, and as we do so, a Hopf bifurcation emerges (Figure A1), suggesting that a stronger feedback induces periodic behavior.

422 Exploring this relationship further, we see that a higher nutrient inflow promotes oscillations, but  
 423 only if the negative feedback is not too strong. As feedback strength increases, less nutrient inflow is  
 424 needed to induce a Hopf bifurcation. This result suggests that EPS can destabilize or stabilize the  
 system depending on how strongly EPS feeds back on nutrient dynamics (Figure A1).



**Figure A1:** (a): Algal bifurcation diagram of feedback strength  $p$  showing that as  $p$  increases a periodic solution appears. (b): EPS bifurcation diagram of feedback strength  $p$  showing that as  $p$  increases a periodic solution appears (c): Two-parameter bifurcation graph. On one side of the curve, there is a stable spiral towards the equilibrium point  $g_1$ , while on the other side, there is a periodic orbit around  $g_1$ . All parameter values are as in Table 1.

425

## 426 Appendix B: Regime 2 trace is always negative

427 We show here that the trace  $Tr(J_{r_1})$  in (12) is negative by proof by contradiction. Assume both  
 428  $b > 0$  and  $Tr(J_{r_1}) > 0$ . Then it follows from (10) that  $\text{sign}(b) = \text{sign}(afe^{-hA^*} - cA^*)$  and the  
 429  $\text{sign}(Tr(J_{r_1})) = \text{sign}(-e^{-hA^*}af^2 + cA^*)$  from equation (12). Using the fact that  $b > 0$  and  $Tr(J_{r_1}) > 0$   
 430 gives:

$$\begin{aligned} afe^{-hA^*} &> cA^* > af^2e^{-hA^*} \\ \implies f &> \frac{cA^*}{ae^{-hA^*}} > f^2 \end{aligned}$$

431 However, assuming that  $f > 1$  implies that  $f^2 > f$ , contradicting the inequality  $f > f^2$ . Thus, there  
 432 is no regime where both  $Tr(J_{r_1}) > 0$  and  $b > 0$ . Thus,  $Tr(J_{r_1}) > 0 \implies b < 0$ , which lies outside  
 433 the biologically meaningful parameter regime.

## 434 Appendix C: Nondimensional parameters for regime 1 and 3

$$\begin{aligned} N^* &= \frac{N}{\gamma} & a &= \frac{\phi}{\gamma\delta} \\ A^* &= \frac{A}{\gamma} & b &= \frac{\psi}{\delta} \\ E^* &= \frac{E}{\mu} & c &= \frac{\nu}{\delta} \\ t &= \frac{\tau}{\eta} & d &= \frac{\rho\gamma}{\mu\eta} \\ f &= \xi c & \epsilon &= \frac{\eta}{\delta} \end{aligned}$$

435 **Appendix D: Nondimensional parameters for regime 2**

$$\begin{array}{ll} N^* & = \frac{N}{\gamma} & a & = \frac{\phi}{\gamma\delta} \\ A^* & = \frac{A}{\gamma} & b & = \frac{\psi}{\delta} \\ t & = \frac{\tau}{\delta} & c & = \frac{\nu}{\delta} \\ f & = \xi c & h & = \frac{\sigma\gamma}{\mu} \end{array}$$

## 436 References

- 437 Arrigo KR, Van Dijken G, Pabi S (2008) Impact of a shrinking Arctic ice cover on marine primary  
 438 production. *Geophysical Research Letters* 35(19):L19603. <https://doi.org/10.1029/2008GL035028>
- 439 Brown TA, Galicia MP, Thiemann GW, et al (2018) High contributions of sea ice derived carbon in  
 440 polar bear (*Ursus maritimus*) tissue. *PLOS ONE* 13(1):e0191631. <https://doi.org/10.1371/journal.pone.0191631>
- 441
- 442 Collins RE, Carpenter SD, Deming JW (2008) Spatial heterogeneity and temporal dynamics of  
 443 particles, bacteria, and pEPS in Arctic winter sea ice. *Journal of Marine Systems* 74(3–4):902–917.  
 444 <https://doi.org/10.1016/j.jmarsys.2007.09.005>
- 445 Ewert M, Deming JW (2013) Sea ice microorganisms: Environmental constraints and extracellular  
 446 responses. *Biology* 2(2):603–628. <https://doi.org/10.3390/biology2020603>
- 447 Gradinger R (2009) Sea-ice algae: Major contributors to primary production and algal biomass in  
 448 the Chukchi and Beaufort Seas during May/June 2002. *Deep Sea Research Part II: Topical Studies*  
 449 *in Oceanography* 56(17):1201–1212. <https://doi.org/10.1016/j.dsr2.2008.10.016>
- 450 Hunke E, Allard R, Bailey DA, et al (2024) CICE-Consortium/Icepack: Icepack 1.5.0. <https://doi.org/10.5281/zenodo.14188409>
- 451
- 452 Huppert A, Blasius B, Stone L (2002) A Model of Phytoplankton Blooms. *The American Naturalist*  
 453 159(2):156–171. <https://doi.org/10.1086/324789>
- 454 Kim K, Ha S, Kim BK, et al (2020) Carbon and nitrogen uptake rates and macromolecular compo-  
 455 sitions of bottom-ice algae and phytoplankton at Cambridge Bay in Dease Strait, Canada. *Annals*  
 456 *of Glaciology* 61(82):106–116. <https://doi.org/10.1017/aog.2020.17>
- 457 Krembs C, Eicken H, Junge K, et al (2002) High concentrations of exopolymeric substances in Arctic  
 458 winter sea ice: Implications for the polar ocean carbon cycle and cryoprotection of diatoms. *Deep-*  
 459 *Sea Research Part I: Oceanographic Research Papers* 49(12):2163–2181. [https://doi.org/10.1016/S0967-0637\(02\)00122-X](https://doi.org/10.1016/S0967-0637(02)00122-X)
- 460
- 461 Krembs C, Eicken H, Deming JW (2011) Exopolymer alteration of physical properties of sea ice  
 462 and implications for ice habitability and biogeochemistry in a warmer Arctic. *Proceedings of the*  
 463 *National Academy of Sciences* 108(9):3653–3658. <https://doi.org/10.1073/pnas.1100701108>
- 464 Lake RA, Lewis EL (1970) Salt rejection by sea ice during growth. *Journal of Geophysical Research*  
 465 75(3):583–597. <https://doi.org/10.1029/JC075i003p00583>
- 466 Lannuzel D, Tedesco L, Van Leeuwe M, et al (2020) The future of Arctic sea-ice biogeochemistry  
 467 and ice-associated ecosystems. *Nature Climate Change* 10(11):983–992. <https://doi.org/10.1038/s41558-020-00940-4>
- 468
- 469 Miller LA, Fripiat F, Else BG, et al (2015) Methods for biogeochemical studies of sea ice: The  
 470 state of the art, caveats, and recommendations. *Elementa: Science of the Anthropocene* 3:000038.  
 471 <https://doi.org/10.12952/journal.elementa.000038>
- 472 Petrich C, Eicken H (2009) Growth, structure and properties of sea ice. In: *Sea Ice*. John Wiley &  
 473 Sons, Ltd, chap 2, p 23–77, <https://doi.org/10.1002/9781444317145.ch2>
- 474 Riedel A, Michel C, Gosselin M (2006) Seasonal study of sea-ice exopolymeric substances on the  
 475 Mackenzie shelf: Implications for transport of sea-ice bacteria and algae. *Aquatic Microbial Ecology*  
 476 45:195–206. <https://doi.org/10.3354/ame045195>
- 477 Roukaerts A, Deman F, Van Der Linden F, et al (2021) The biogeochemical role of a microbial  
 478 biofilm in sea ice. *Elem Sci Anth* 9(1):001–14. <https://doi.org/10.1525/elementa.2020.00134>
- 479 Steffen KR, Epshteyn Y, Zhu J, et al (2018) Network Modeling of Fluid Transport Through Sea  
 480 Ice with Entrained Exopolymeric Substances. *Multiscale Modeling & Simulation* 16(1):106–124.  
 481 <https://doi.org/10.1137/17M1117513>
- 482 Stoecker DK, Gustafson DE, Baier CT, et al (2000) Primary production in the upper sea ice. *Aquatic*  
 483 *Microbial Ecology* 21(3):275–287. <https://doi.org/10.3354/ame021275>
- 484 Tedesco L, Vichi M (2014) Sea ice biogeochemistry: A guide for modellers. *PLoS ONE* 9(2):e89217.  
 485 <https://doi.org/10.1371/journal.pone.0089217>
- 486 Thomas DN, Dieckmann GS (2010) *Sea Ice*, 2nd edn. Wiley-Blackwell, <https://doi.org/10.1002/9781444317145>
- 487
- 488 Zong Z (2022) A Random Pore Model of sea ice for predicting its mechanical properties. *Cold Regions*  
 489 *Science and Technology* 195:103473. <https://doi.org/10.1016/j.coldregions.2021.103473>